

## 13 — The Gray & Eddington Approximations [*Revision* : 1.3]

- Gray atmospheres

- Plane-parallel RTE still difficult to solve because vertical optical depth is wavelength dependent
- Useful simplification of **gray atmosphere**: assume opacity is independent of wavelength:  $\kappa_\lambda = \bar{\kappa}$ . Then,

$$\tau_v(z) = \int_z^0 \bar{\kappa}(z) \rho(z) dz$$

and

$$\cos \theta \frac{dI}{d\tau_v} = I - S$$

where

$$I = \int_0^\infty I_\lambda d\lambda, \quad S = \int_0^\infty S_\lambda d\lambda$$

- Moments of the radiation field

- Grey, plane-parallel RTE still depends on direction  $\mu$
- To further simplify, integrate RTE over all solid angles:

$$\frac{d}{d\tau_v} \int I \mu d\Omega = \int I d\Omega - S \int d\Omega$$

( $S$  comes out from under integral sign, because it doesn't depend on angle)

- Recalling from notes 9, mean intensity:

$$\langle I \rangle = \frac{1}{4\pi} \int I d\Omega = \frac{1}{2} \int_{-1}^1 I d\mu,$$

and flux:

$$F = \int I \mu d\Omega = 2\pi \int_{-1}^1 I \mu d\mu$$

- So,

$$\frac{dF}{d\tau_v} = 4\pi(\langle I \rangle - S)$$

- Another equation is obtained by multiplying RTE by  $\mu$ , and integrating over all solid angles:

$$\frac{d}{d\tau_v} \int I \mu^2 d\Omega = \int I \mu d\Omega - S \int \mu d\Omega$$

- Recalling from notes 9, radiation pressure:

$$P_{\text{rad}} = \frac{1}{c} \int I \mu^2 d\Omega = \frac{2\pi}{c} \int_{-1}^1 I \mu^2 d\mu$$

and so:

$$\frac{dP_{\text{rad}}}{d\tau_v} = \frac{F}{c}$$

- Equations for  $dF/d\tau_v$  and  $dP_{\text{rad}}/d\tau_v$  are **moment equations**;  $n^{\text{th}}$  moment of radiation field defined by

$$M_n = \int I \mu^n d\Omega$$

- Radiative equilibrium

- In **radiative equilibrium**, no net energy added or subtracted as radiation travels up through atmosphere — amount absorbed = amount emitted:

$$\kappa \langle I \rangle = \kappa S$$

- From first moment equation, this means that

$$\frac{dF}{d\tau} = 4\pi(\langle I \rangle - S) = 0$$

i.e., flux  $F$  is constant throughout atmosphere

- From second moment equation, with  $F$  constant:

$$P_{\text{rad}} = \frac{1}{c} F \tau_v + C$$

where  $C$  is constant of integration

- **However**, how do we find out  $P_{\text{rad}}$ ? Could take next higher-order moment of RTE, but that will just introduce an additional term  $M_3$  — and so on, *ad infinitum*
- In fact, we need **closure relation**: independent piece of information that will relate  $P_{\text{rad}}$  to  $\langle I \rangle$

- Eddington approximation

- To find closure relation, make simplifying **Eddington approximation** about radiation field: specific intensity is constant over upward and downward hemispheres:

$$I(\mu) = \begin{cases} I_{\text{out}} & \mu > 0 \\ I_{\text{in}} & \mu < 0 \end{cases}$$

- Substitute into definitions above for various moments gives

$$\langle I \rangle = \frac{1}{2}(I_{\text{out}} + I_{\text{in}})$$

$$F = \pi(I_{\text{out}} - I_{\text{in}})$$

$$P_{\text{rad}} = \frac{2\pi}{3c}(I_{\text{out}} + I_{\text{in}}) = \frac{4\pi}{3c}\langle I \rangle$$

- Last equation gives the desired closure relation — an independent piece of information relating radiation pressure to mean intensity
- Combining with earlier equation for radiation pressure:

$$\frac{4\pi}{3c}\langle I \rangle = \frac{1}{c} F \tau_v + C$$

- Eliminate  $C$  by requiring that  $I_{\text{in}} = 0$  at  $\tau_v = 0$ :

$$\frac{4\pi}{3}\langle I \rangle = F \left( \tau_v + \frac{2}{3} \right)$$

- Final step: since  $F$  is constant in atmosphere, and at outer layers  $F = \sigma T_{\text{eff}}^4$  (**definition** of effective temperature!),

$$\langle I \rangle = \frac{3\sigma}{4\pi} T_{\text{eff}}^4 \left( \tau_v + \frac{2}{3} \right)$$

This then gives source function throughout atmosphere, and hence full radiation field via formal solution

- If atmosphere is in **local thermodynamic equilibrium**,  $S = B = \sigma T^4/\pi$ , and so we have **temperature structure**:

$$T^4 = \frac{3}{4}T_{\text{eff}}^4 \left( \tau_v + \frac{2}{3} \right)$$